



Fairmode Technical Meeting
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Validation of Complex Data Assimilation Methods

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Forschungszentrum Jülich

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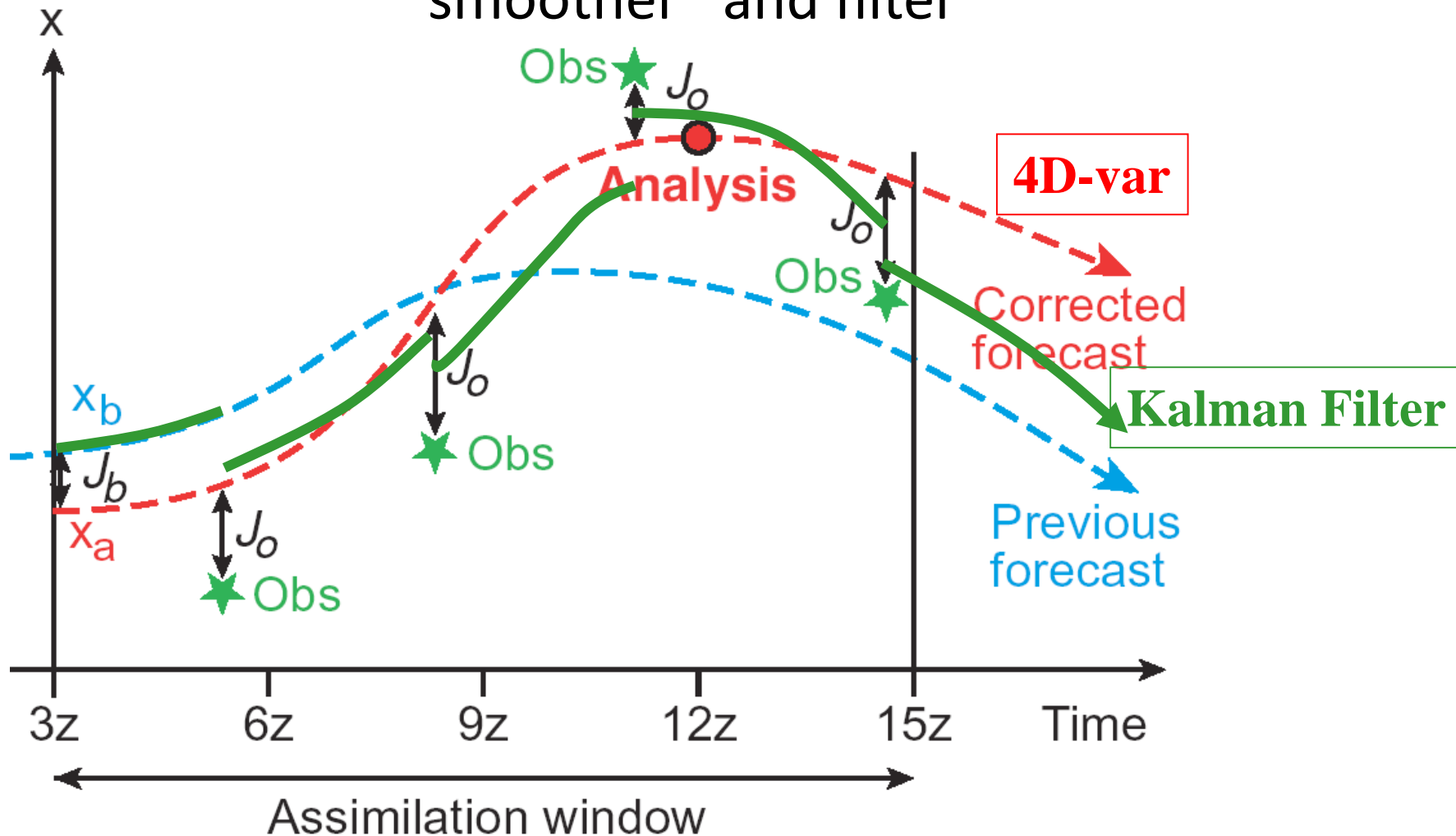
1. Intro.: What are *complex data assimilation methods*?
2. Observability: Do observations sustain assimilation results?
3. Practical verification: Validation by forecast skills
4. A posteriori Validation: Is the analysis consistent?

What are *complex data assimilation methods*?

→ spatio-temporal techniques

2 types of assimilation algorithms:

“smoother” and filter



The 4-dimensional variational technique: *Optimize over an assimilation window, then forecast*

Emission Rate Optimization

minimize cost function

$$J(\mathbf{x}(t_0), \mathbf{e}) = \frac{1}{2}(\mathbf{x}^b(t_0) - \mathbf{x}(t_0))^T \mathbf{B}_0^{-1}(\mathbf{x}^b(t_0) - \mathbf{x}(t_0)) + \frac{1}{2} \int_{t_0}^{t_N} (\mathbf{e}_b(t) - \mathbf{e}(t))^T \mathbf{K}^{-1}(\mathbf{e}_b(t) - \mathbf{e}(t)) dt + \frac{1}{2} \int_{t_0}^{t_N} (\mathbf{y}^0(t) - H[\mathbf{x}(t)])^T \mathbf{R}^{-1}(\mathbf{y}^0(t) - H[\mathbf{x}(t)]) dt$$

deviations from background initial state

deviations from a priori emission rates

model deviations from observations

- $\mathbf{x}^b(t_0)$ background state at $t = 0$
- $\mathbf{x}(t)$ model state at time t
- $\mathbf{e}_b(t_0)$ background emission rate at $t = 0$
- $\mathbf{e}(t)$ emission rate field at time t
- \mathbf{K} emission rate error covariance matrix
- $H[]$ forward interpolator
- $\mathbf{y}^0(t)$ observation at time t
- \mathbf{B}_0 background error covariance matrix

Kalman filter: basic equations

Forecast steps:

a) the atmospheric state

$$\mathbf{x}^f(t_i) = \mathbf{M}(t_i, t_{i-1})\mathbf{x}^a(t_{i-1}) + \boldsymbol{\eta}$$

b) the forecast error covariance matrix

$$\mathbf{P}_i^b = \mathbf{M}(t_i, t_{i-1})\mathbf{P}_{i-1}^a\mathbf{M}^T(t_i, t_{i-1}) + \mathbf{Q}$$

Analysis steps:

a) the atmospheric state

$$\mathbf{x}^a(t_i) = \mathbf{x}^b(t_i) + \mathbf{K}_i\mathbf{d}_i, \quad (1)$$

$$\mathbf{K}_i := \mathbf{P}_i^b\mathbf{H}_i^T(\mathbf{H}_i\mathbf{P}_i^b\mathbf{H}_i^T + \mathbf{R}_i)^{-1} \in \mathcal{R}^{n \times p_i} \quad (2)$$

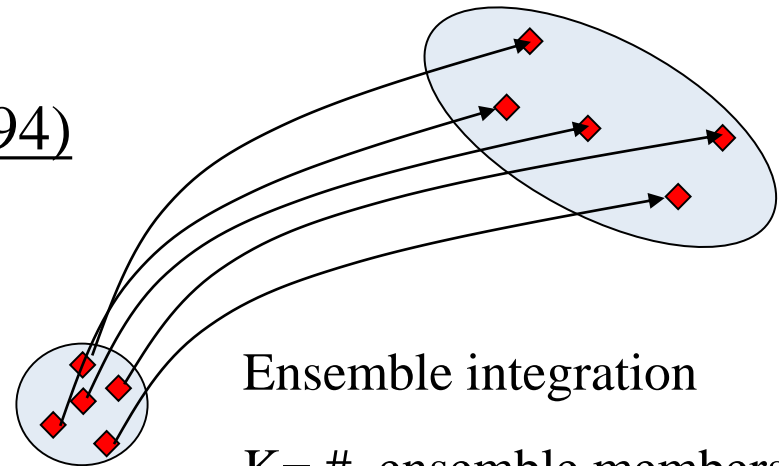
and b) the analysis error covariance matrix

$$\mathbf{P}_i^a = (\mathbf{I} - \mathbf{K}_i\mathbf{H}_i)\mathbf{P}_i^b. \quad (3)$$

Computational challenge: Background Error Covariance Matrix \mathbf{P}^b

1. Ensemble approach: (e.g. Evensen, 1994)

$$B_{ij} = \frac{1}{K} \sum_{n=1}^K (x_i^n - \bar{x}_i)(x_j^n - \bar{x}_j)$$



Ensemble integration

$K = \#$ ensemble members

i, j grid cells

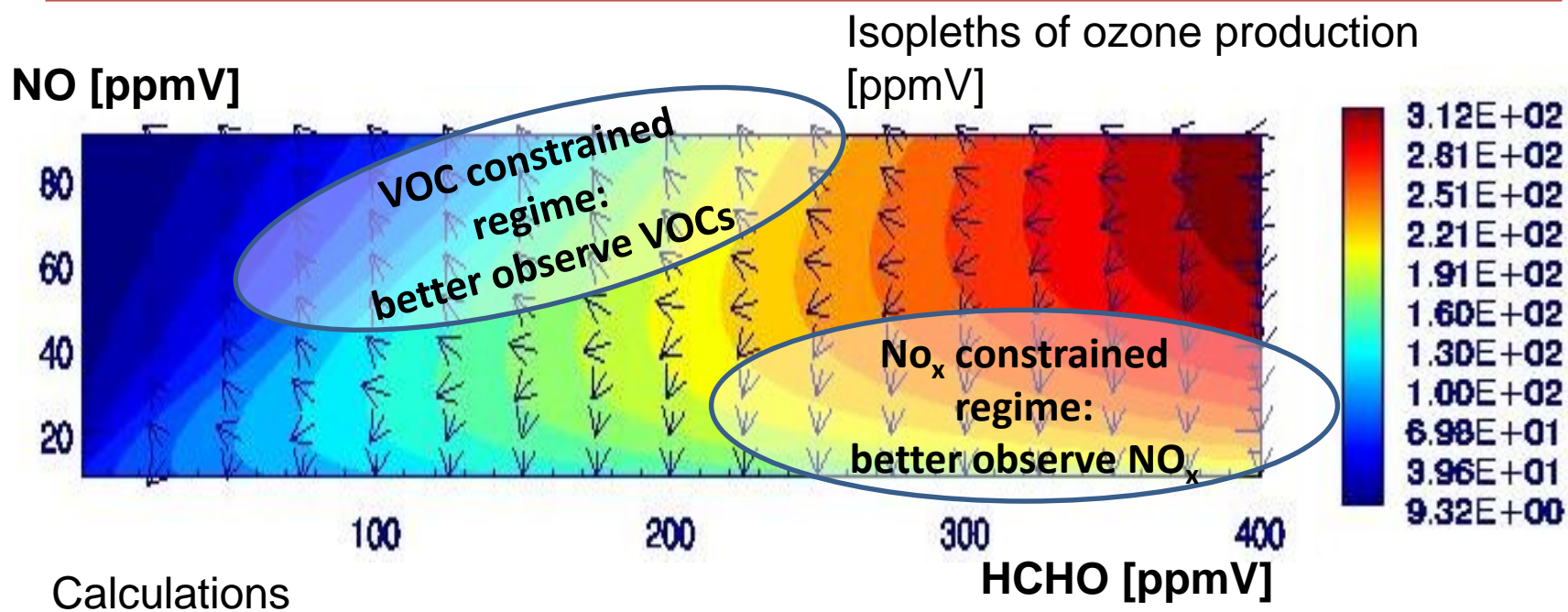
2. Observability: Do observations sustain assimilation results?

Observation network design

Is the forecasted system sensitive to available observations?

- Observation System Simulation Experiments (OSSEs)
- Targeted observations

Is NO_x the key to ozone production? And consequently, its observation [^]the key to better forecast?



- ✓ within a fixed time span
- ✓ initial concentrations of NO / HCHO were varied
- ✓ change of final concentration is given by colour
- ✓ gradients (SVs) of maximyl ozone production given by arrows

How can we optimize the observation configuration?

Given CTM (here RACM and EURAD-IM) acting as tan.-lin. model operator \mathcal{L} :

$$\delta \mathbf{c}(t_F) = \mathcal{L}_{t_I, t_F} \delta \mathbf{c}(t_I), \quad \mathcal{L}_{t_I, t_F} = \left. \frac{\partial \mathcal{M}_{t_I, t_F}}{\partial \mathbf{c}} \right|_{\mathbf{c}(t_I)}$$

1. Berliner et al., (1998) Statistical design:
 “Minimize” the analysis error covariance matrix \mathbf{A} (say, via trace):

$$\min_{\mathbf{H}} \mathbf{A} = \mathbf{B} - \underbrace{\mathbf{B} \mathbf{H}^T (\mathbf{H} \mathbf{B} \mathbf{H}^T + \mathbf{R})^{-1} \mathbf{H} \mathbf{B}}_{\text{to be maximized by } \mathbf{H}}$$

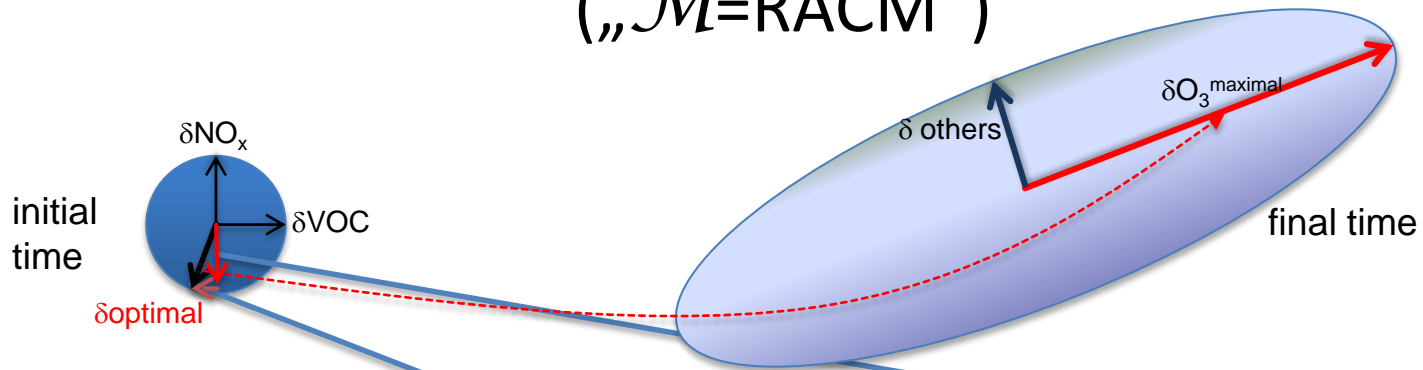
For this find maximal eigenvectors as observation operators \mathbf{H} , which configure observations.

$$\mathcal{L}_{t_I, t_F} \mathbf{B} \mathcal{L}_{t_I, t_F}^T \mathbf{H}^T = \lambda \mathbf{H}^T$$

2. Palmer (1995) Singular vector analysis:
 Observe maximal SV configuration:

$$\max_{\delta \mathbf{c}(t_I)} \frac{\|\delta \mathbf{c}(t_F)\|_{\mathbf{B}}^2}{\|\delta \mathbf{c}(t_I)\|_{\mathbf{B}}^2} = \max_{\delta \mathbf{c}(t_I)} \frac{\delta \mathbf{c}(t_I)^T \mathcal{L}_{t_I, t_F}^T \mathbf{B} \mathcal{L}_{t_I, t_F} \delta \mathbf{c}(t_I)}{\delta \mathbf{c}(t_I)^T \mathbf{B} \delta \mathbf{c}(t_I)},$$

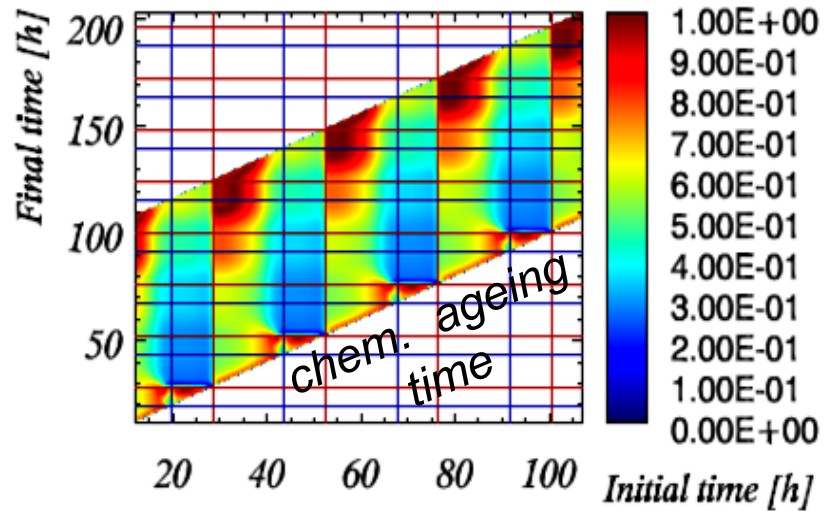
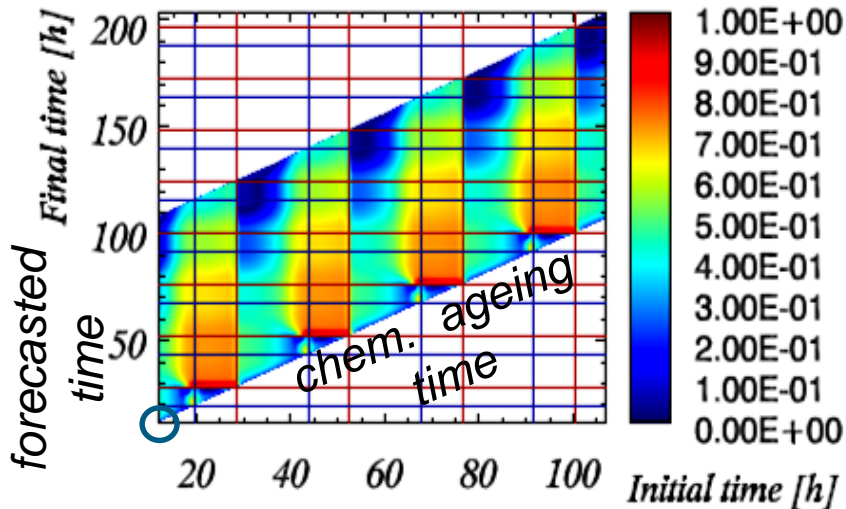
Basic 0-D Regional Atm. Chemistry Mechanism („ \mathcal{M} =RACM“)



- Optimal perturbations (Singular Vectors) for scenario MARINE**

1st Grouped Singular Vectors (δ VOC)

1st Grouped Singular Vectors (δ NO_x)

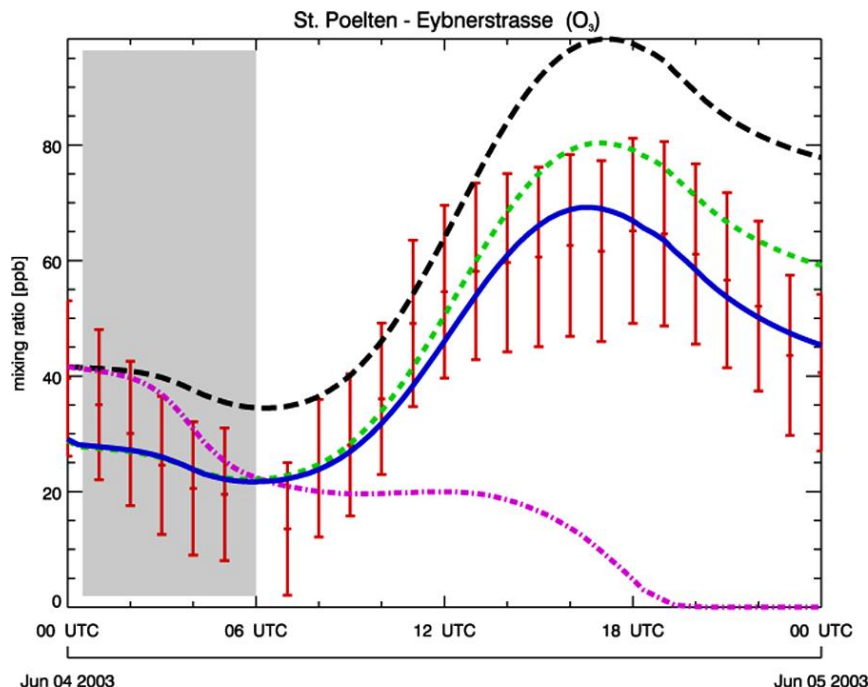


not | very
important to observe

— sunrise — sunset

3. Practical verification: Validation by forecasts

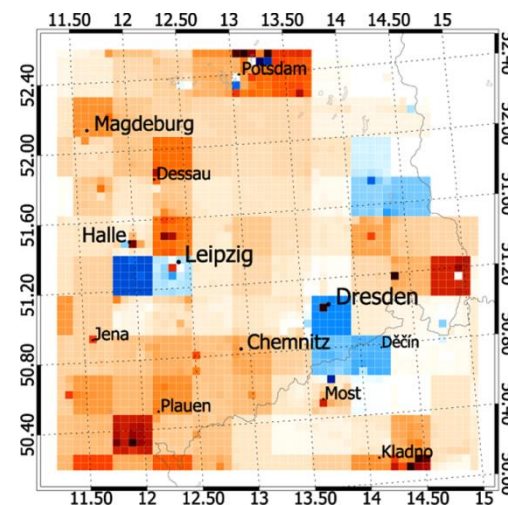
Analysis of emissions by 4D-var (VERTIKO)



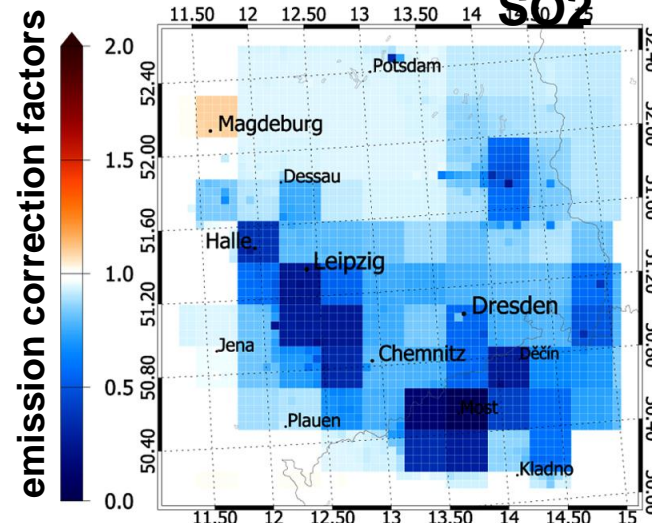
Observed and analysed ozone evolution at St. Poelten Vertical bars: ozone observations with error estimates.

- Control run without data assimilation.
- initial value optimisation.
- - - - - emission factor optimisation.
- joint initial value and emission factor optimisation (Strunk et al., 2011)

NO2



SO2

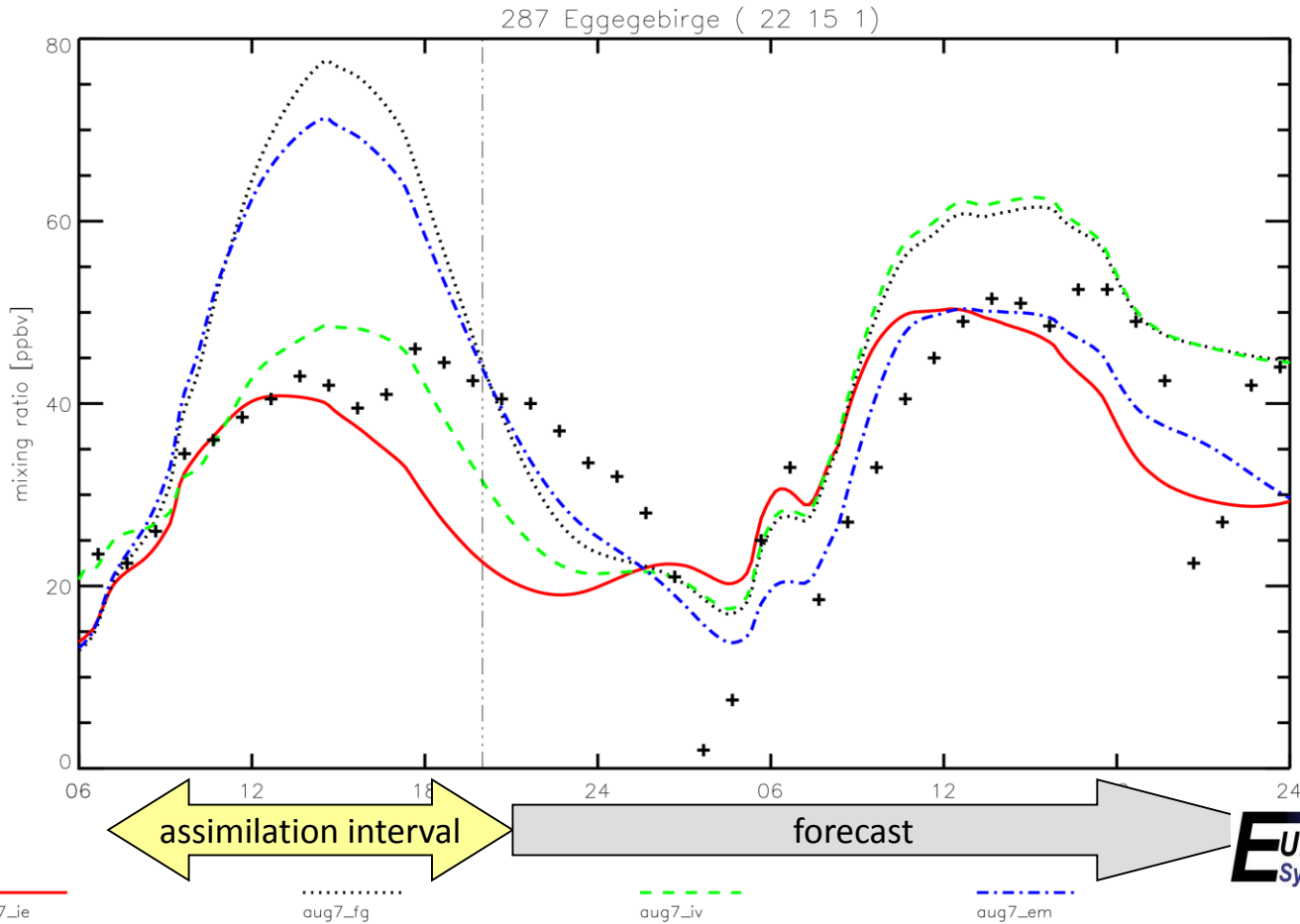


4. Focus: joint emission rate initial value optimisation

Semi-rural measurement site **Eggegebirge**

7. August

8. August 1997



+ observations
no optimisation

initial value opt.

emis. rate opt.

joint emis +
ini val opt.



aug7_ie

aug7_fg

aug7_iv

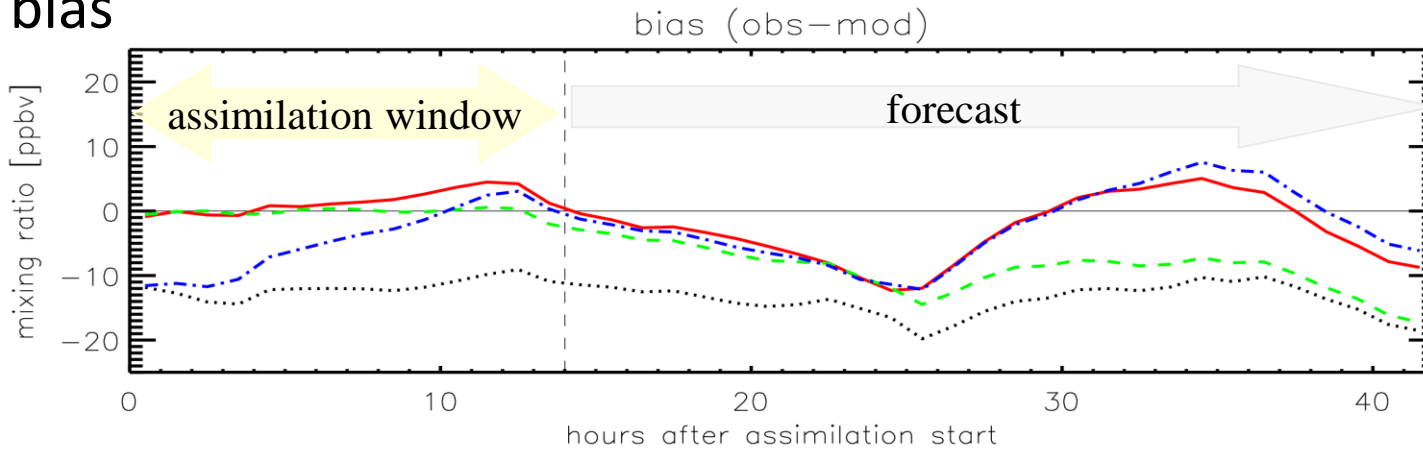
aug7_em

How long does data assimilation have an impact?

Answer gas phase

12-24 hours, dependent on optimisation

bias

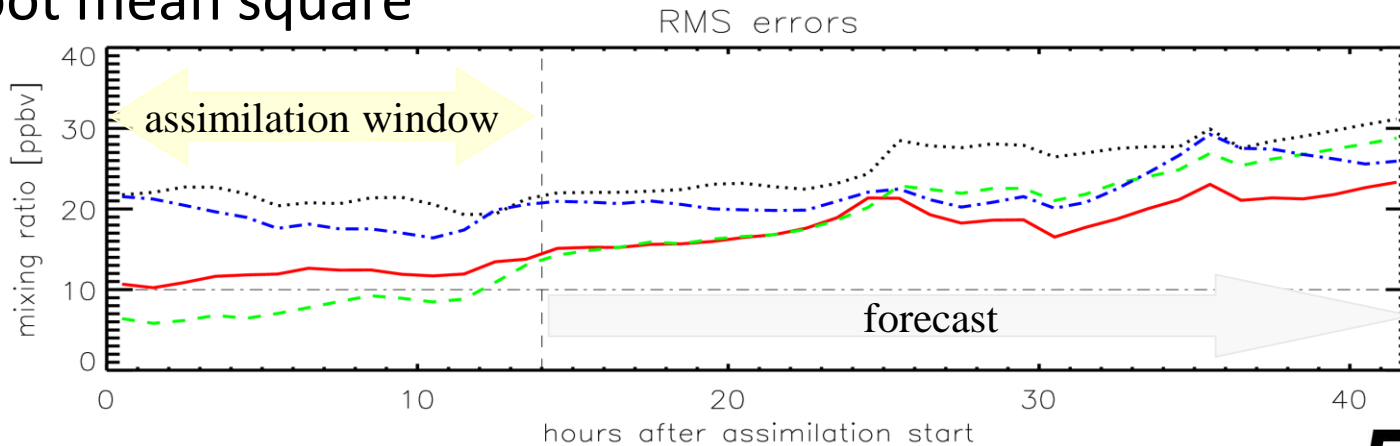


+ observations
no optimisation

initial value opt.

emis. rate opt.

root mean square



joint emis +
ini val opt.

aug7_ie

aug7_fg

aug7_iv

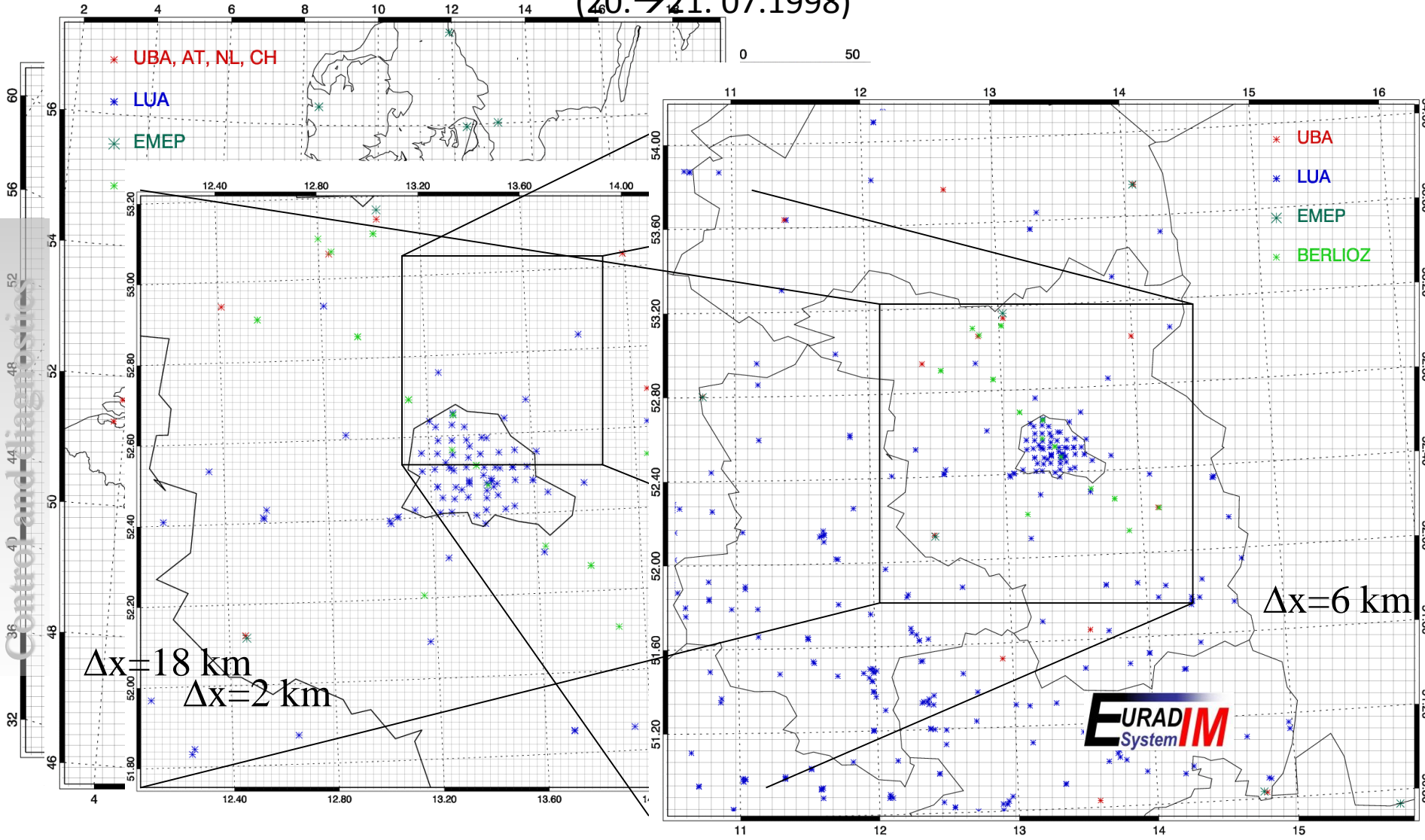
aug7_em



Which is the requested resolution?

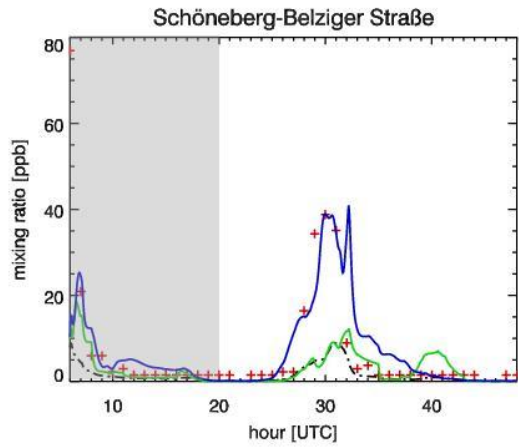
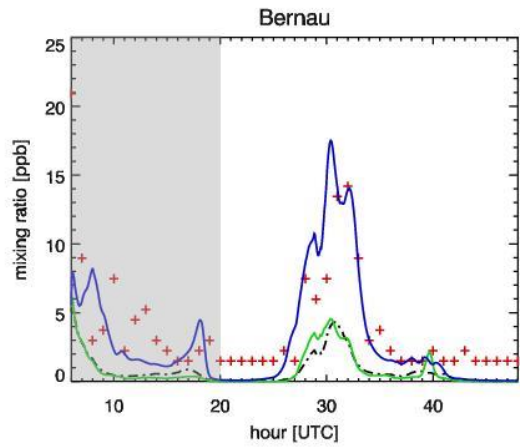
BERLIOZ grid designs and observational sites

(20. → 21. 07.1998)



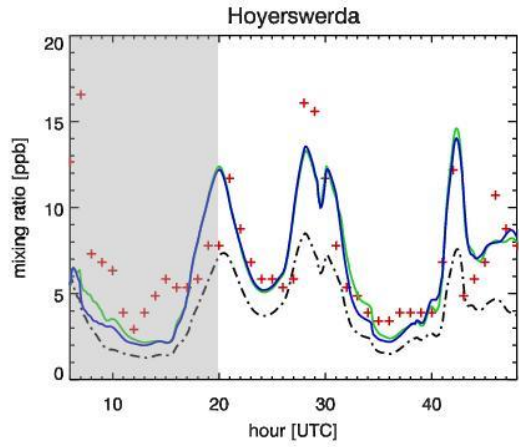
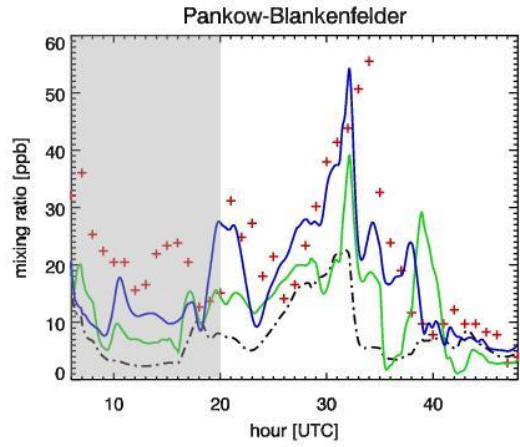
Some BERLIOZ examples of NOx assimilation (20. → 21. 07.1998)

NO



Time series for selected NOx stations on nest 2.
+ observations,
- - - no assimilation,
- N1 assimilation (18 km),
- N2 assimilation (6 km),
- grey shading: assimilated observations, others forecasted.

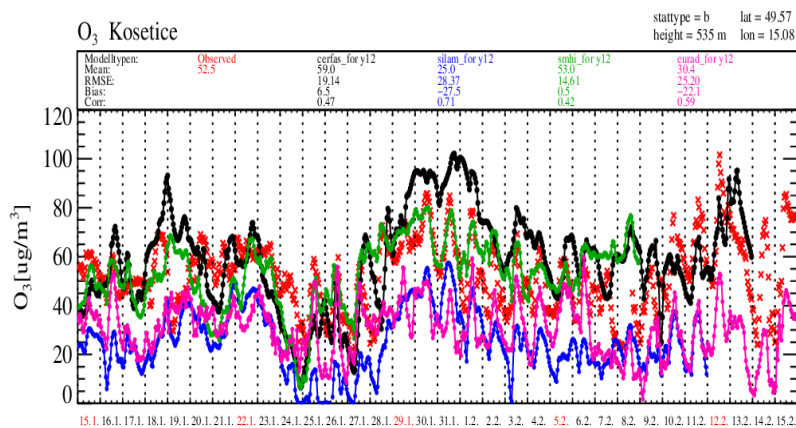
NO₂



Validation by measurements withheld

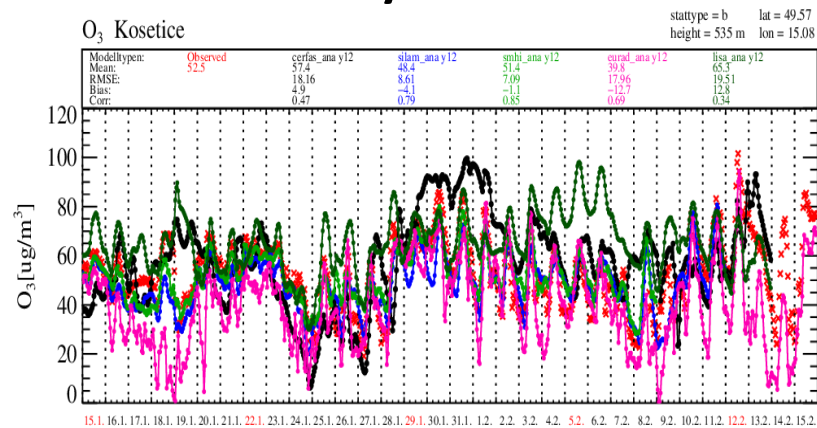
(extract from MACC III EDA report draft)

Forecast

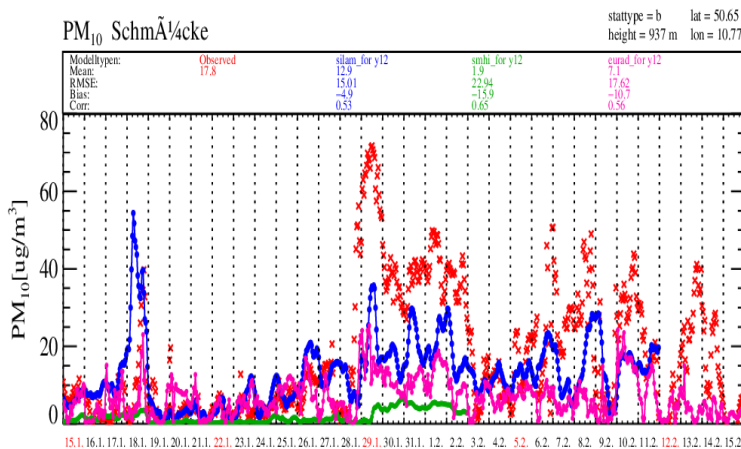


15. Jan., 0 UTC – 15. Feb. 2012, 23 UTC

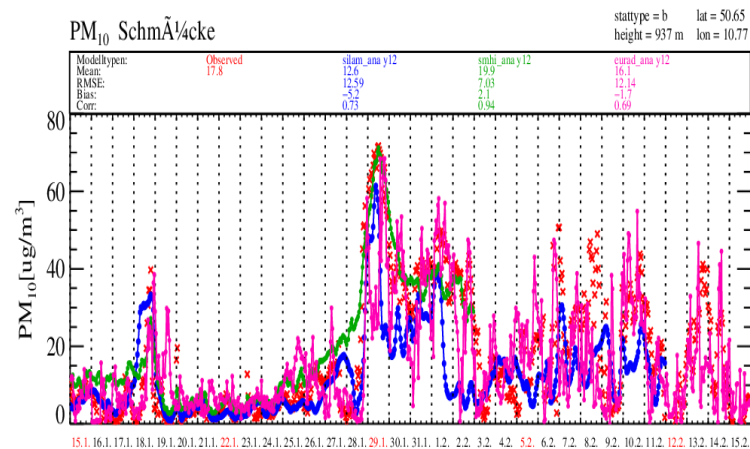
Analyses



15. Jan., 0 UTC – 15. Feb. 2012, 23 UTC



15. Jan., 0 UTC – 15. Feb. 2012, 23 UTC

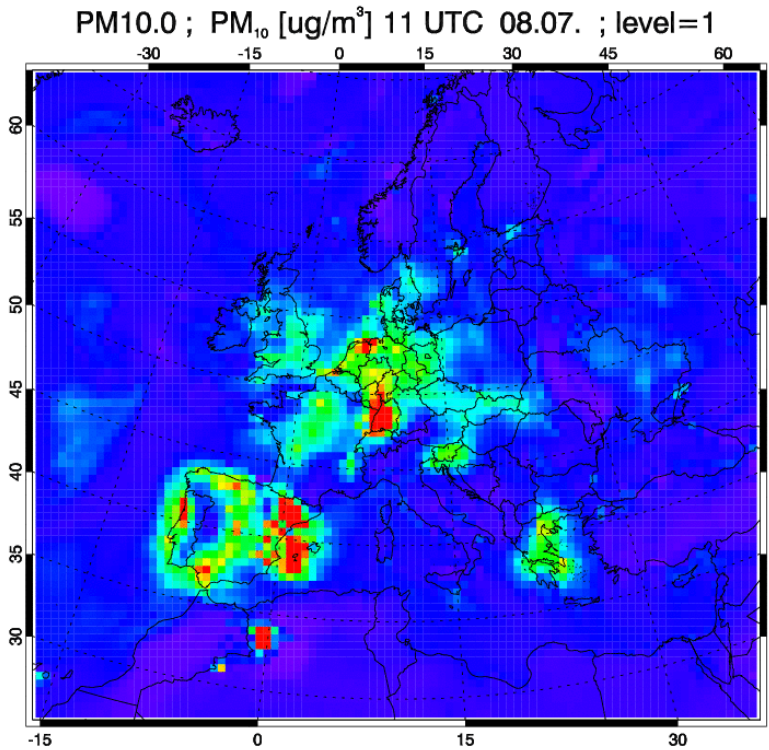


15. Jan., 0 UTC – 15. Feb. 2012, 23 UTC

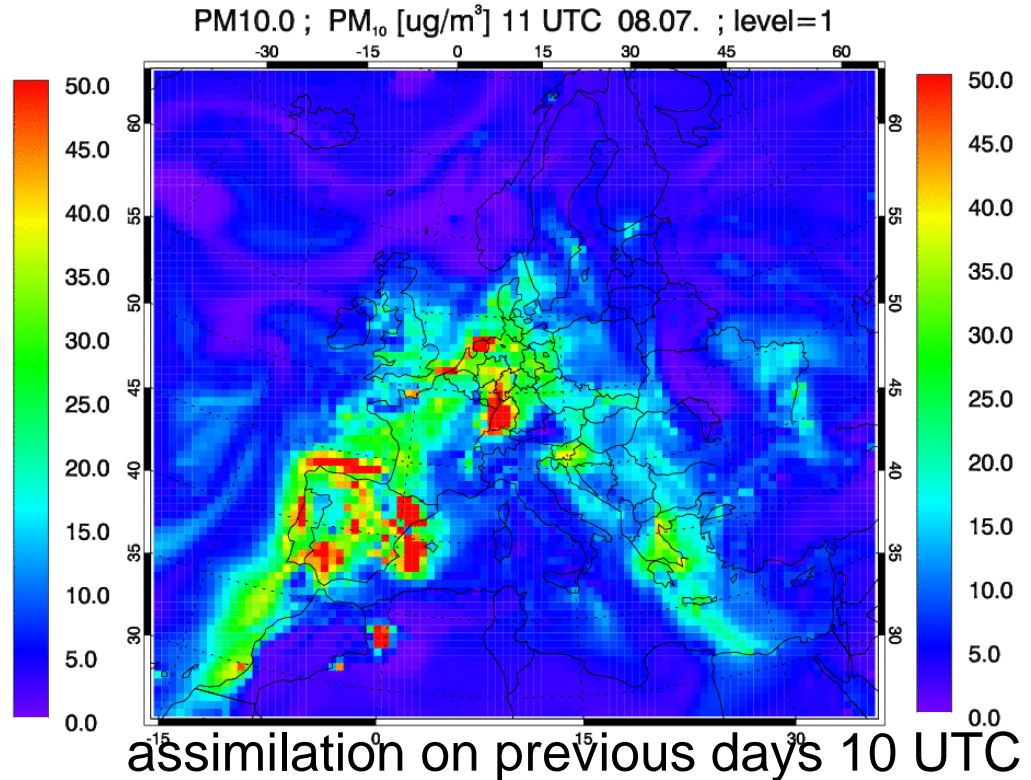
How long does data assimilation have an impact?

Answer aerosol phase

aerosol data assimilation effects accumulate



No previous assimilation
only 1 day: 14. July 2003

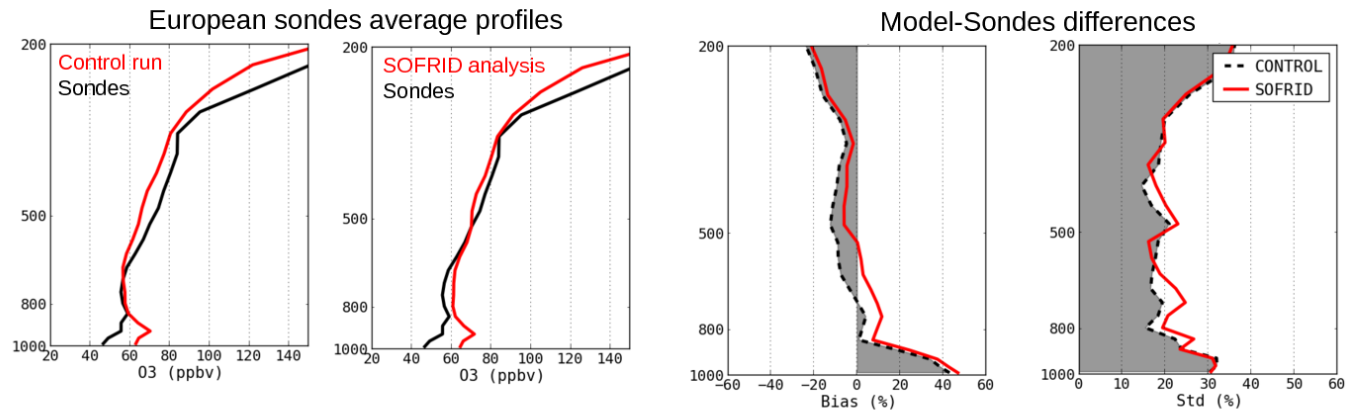


Accumulation of retrieval information over
14 days

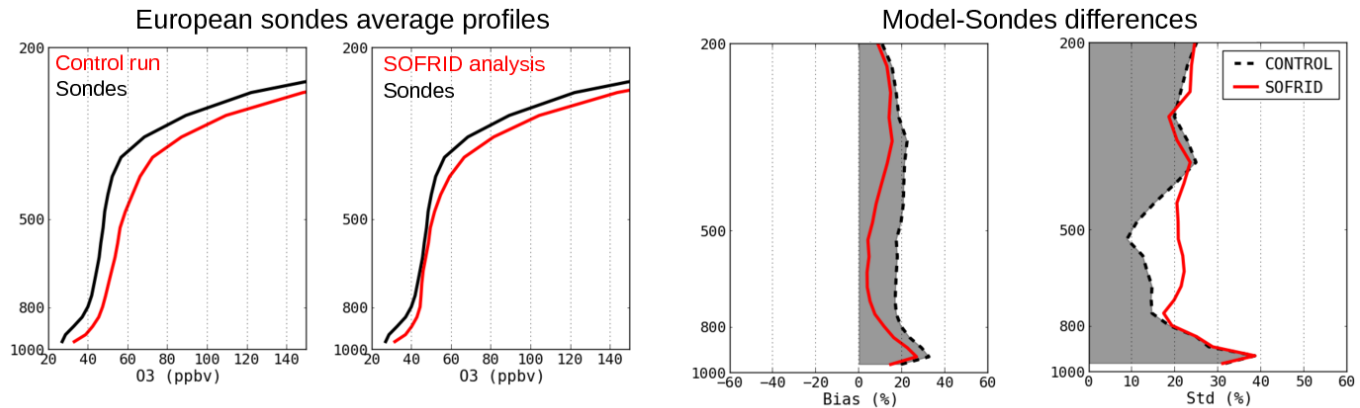
MOCAGE satellite data assimilation: IASI SOFRID O₃ re-analysis (CERFACS)

Validation of IASI analysis with ozonesonde data:
BIAS = model minus observations

O₃ profiles in July 2010:



O₃ profiles in Jan 2012:



- Bias reduced in the free troposphere
- Surface ozone impact is minor
- MOZAIC-IAGOS as additional validation? (only 2012 available)

Courtesy E. Emili, CERFACS

4. A posteriori Validation: Is the analysis consistent?

a posteriori validation of data assimilation results

Assumptions:

- Gaussian error distribution assumption sufficiently valid
- First guess not too far from “solution” (tangent-linear approximation must hold)
- A priori defined error covariances (background, observations)

Necessary condition
for a posteriori
validation:
adjust B and R such
that:

at the minimum:

$$J_{min} = 1/2 d^T (\mathbf{HBH}^T + \mathbf{R})^{-1} d$$

$$d := y - Hx^a$$

p number of observations

Expectation
Variance

$$\mathcal{E}[J_{min}] = p/2$$

$$\mathcal{V}[J_{min}] = p/2$$

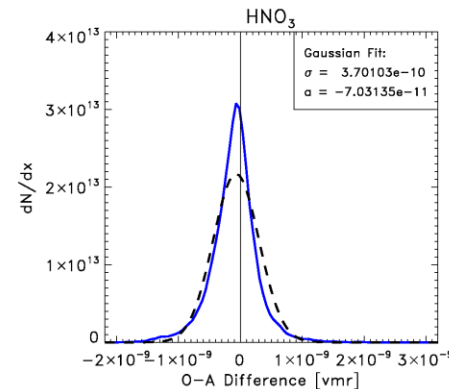
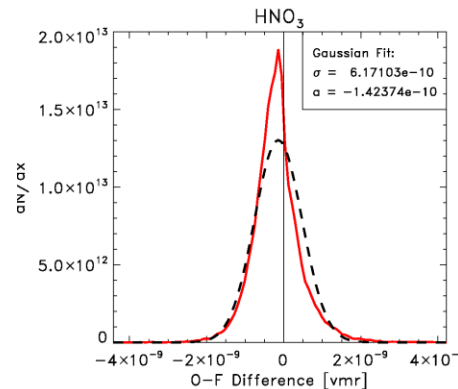
Evaluating the Gaussian error distribution assumption

SACADA

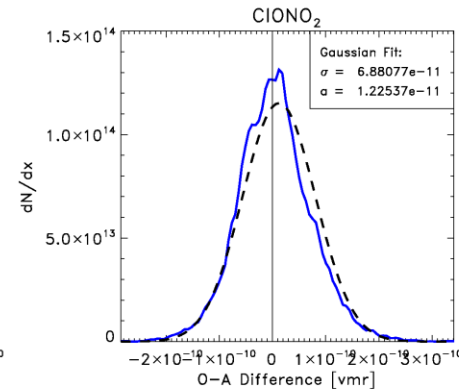
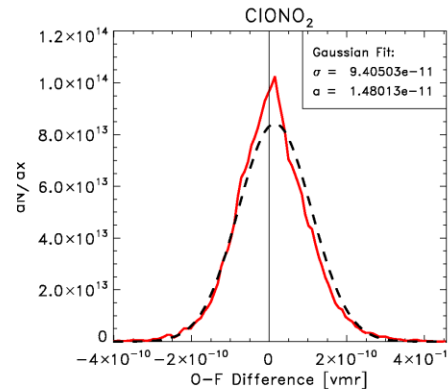
O-F differences (left column) and

O-A differences (right column)

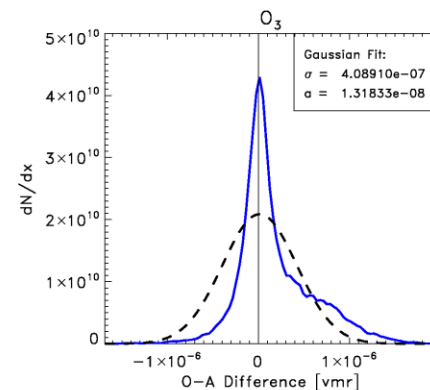
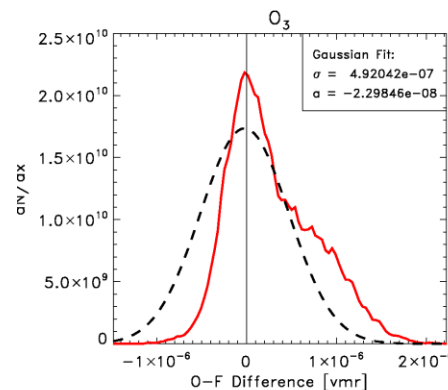
Dotted line represents a Gaussian with same variance as the data



HNO₃

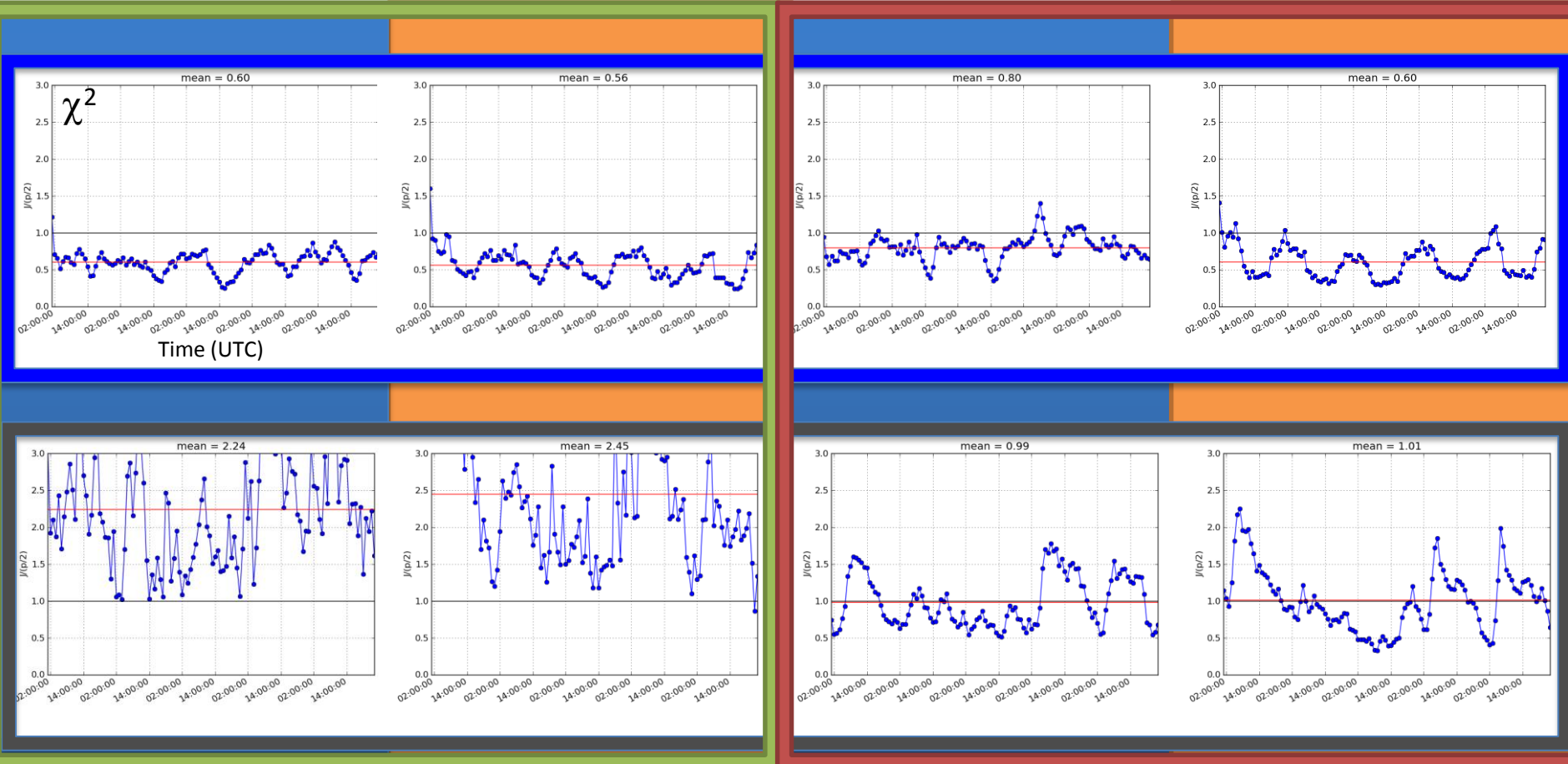
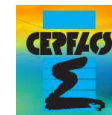


ClONO₂



O₃

χ^2 validation MOCAGE



- Surface O₃ assimilated
- Surface NO₂ assimilated
- Winter period (1-2-2008, 6-8,2008)
- Summer period (1-8-2008, 6-8-2008)
- Only rural background sites assimilated
- Only urban background sites assimilated

Comments:

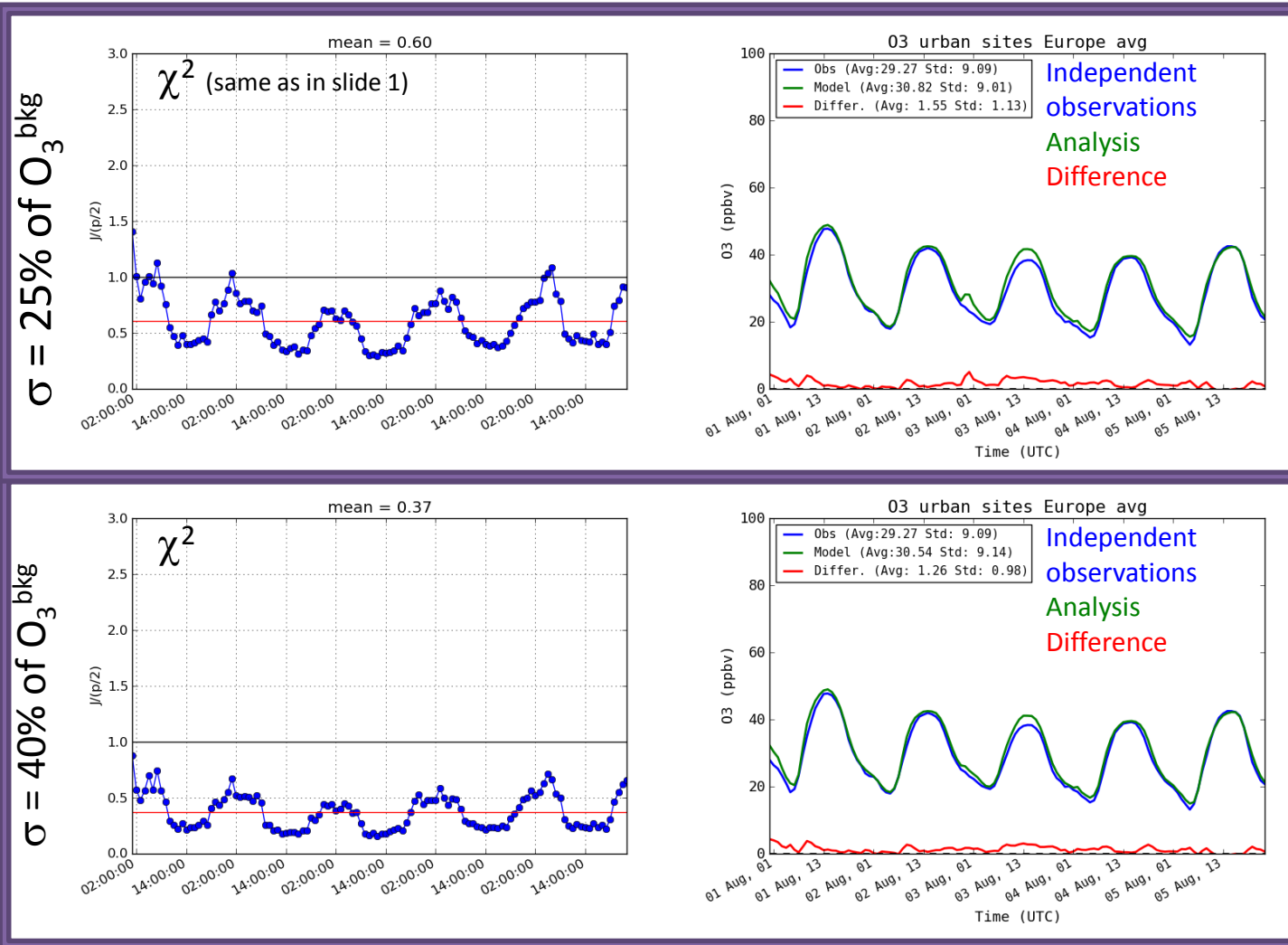
- O₃
 - the urban case is the only case with a distinct winter-summer behavior (higher χ^2 in winter)
 - presence of diurnal variability in all cases
- NO₂
 - large differences between rural/urban cases
 - strong variations in the rural case
 - presence of diurnal variability in all cases
 - no evidence of significant seasonality

Courtesy E. Emili, CERFACS

χ^2 validation MOCAGE

What is the impact of a low χ^2 in terms of validation with an independent dataset?

Example: O₃ background urban sites assimilated in summer, validation against sites kept out from the assimilation, two choices of the background error variance σ



Comments:
 Case 2 ($\sigma = 40\%$) has lower χ^2 but better analysis scores. A better χ^2 does not always imply a better analysis, because χ^2 stats do not consider model biases.

Conclusions

- Atmospheric chemistry is a highly coupled nonlinear dynamic system, which is best addressed by spatio-temporal data assimilation
- the system must be observed with respect to its sensitivity (NO_x-VOX interaction)
- Forecasts must be shown to improve
- the assimilation result must be consistent: proper balance between a priori and a posteriori knowledge (χ^2 -validation)

Additional illustrations

2. Focus: Can we identify flaws?

A posteriori evaluation

1. χ^2 – validation
2. a posteriori validation in observation space

Theoretical background on a posteriori evaluation

$$J_{\min} = \frac{1}{2} \mathbf{d}^T \mathbf{E} \tilde{\mathbf{d}} \mathbf{d}^T \mathbf{d}$$

$$E(J_{\min}) = \frac{p}{2}$$

Aposteriori validation in observation space

Extended Kalman filter equations

Forecast step: $\mathbf{x}^b(t_i) = M_{i-1} [\mathbf{x}^a(t_{i-1})]$

$$\mathbf{B}(t_i) = \mathbf{L}_{i-1} \mathbf{A}(t_{i-1}) \mathbf{L}_{i-1}^T + \mathbf{Q}(t_{i-1})$$

Analysis step: $\mathbf{x}^a(t_i) = \mathbf{x}^b(t_i) + \mathbf{K}(t_i) (\mathbf{y} - H [\mathbf{x}^b(t_i)])$

$$\mathbf{A}(t_i) = (\mathbf{I} - \mathbf{K}(t_i) \mathbf{H}) \mathbf{B}(t_i)$$

where

M_i := Model operator

L_i := Tangent linear model operator

optimize R and B
directly,
and A indirectly

$$\mathbf{K}(t_i) := \mathbf{B}(t_i) \mathbf{H}^T [\mathbf{R} + \mathbf{H} \mathbf{B}(t_i) \mathbf{H}^T]^{-1}$$

$\mathbf{Q}(t_i)$:= Model error covariance matrix

$\mathbf{B}(t_i)$:= Background error covariance matrix

$\mathbf{A}(t_i)$:= Analysis error covariance matrix

$\mathbf{R}(t_i)$:= Observation error covariance matrix

2. Focus: a posteriori validation

Diagnosis and Tuning of Error Covariances

(Desroziers et al. 2005)

makes the difference

$$E \left\{ \mathbf{d}_b^a \mathbf{d}_b^{oT} \right\} = \mathbf{H} \tilde{\mathbf{B}} \mathbf{H}^T$$
$$E \left\{ \mathbf{d}_a^o \mathbf{d}_b^{oT} \right\} = \tilde{\mathbf{R}}$$

$$\mathbf{d}_b^a := H(\mathbf{x}^a) - H(\mathbf{x}^b)$$

$$\mathbf{d}_a^o := \mathbf{y} - H(\mathbf{x}^a)$$

$$\mathbf{d}_b^o := \mathbf{y} - H(\mathbf{x}^b)$$

If \mathbf{B} and \mathbf{R} are consistently specified, then $\mathbf{B} = \tilde{\mathbf{B}}$ and $\mathbf{R} = \tilde{\mathbf{R}}$ and

$$E \left\{ \mathbf{d}_b^a \mathbf{d}_a^{oT} \right\} = \mathbf{H} \mathbf{A} \mathbf{H}^T$$

**Only a necessary, but not a sufficient condition is fulfilled:
no unique solution**

Tuning of Error Covariances in observation space

(Desroziers et al. 2005)

$$E \left\{ \mathbf{d}_b^a \mathbf{d}_b^{oT} \right\} = \mathbf{HBH}^T \quad (1)$$

$$E \left\{ \mathbf{d}_a^o \mathbf{d}_b^{oT} \right\} = \mathbf{R} \quad (2)$$

$$E \left\{ \mathbf{d}_b^o \mathbf{d}_b^{oT} \right\} = \mathbf{HBH}^T + \mathbf{R} \quad (3)$$

$$E \left\{ \mathbf{d}_b^a \mathbf{d}_a^{oT} \right\} = \mathbf{HAH}^T \quad (4)$$

if \mathbf{B} and \mathbf{R} are correctly specified.

$$\mathbf{d}_b^a := H(\mathbf{x}^a) - H(\mathbf{x}^b)$$

$$\mathbf{d}_a^o := \mathbf{y} - H(\mathbf{x}^a)$$

$$\mathbf{d}_b^o := \mathbf{y} - H(\mathbf{x}^b)$$

in practice: Iterative approach

Practical estimate of diagonal elements of R and B

$$\begin{aligned}(\tilde{\sigma}_i^b)^2 &= (\mathbf{d}_b^a)_i^T (\mathbf{d}_b^o)_i = \sum_{j=1}^{p_i} (\mathbf{y}_j^a - \mathbf{y}_j^b)(\mathbf{y}_j^o - \mathbf{y}_j^b)/p_i \\(\tilde{\sigma}_i^o)^2 &= (\mathbf{d}_a^o)_i^T (\mathbf{d}_b^o)_i = \sum_{j=1}^{p_i} (\mathbf{y}_j^o - \mathbf{y}_j^a)(\mathbf{y}_j^o - \mathbf{y}_j^b)/p_i\end{aligned}$$

Estimate of off-diagonal elements of B

$$(\tilde{\sigma}_{ij}^b)^2 = \sum_{\substack{i,j=1 \\ i \neq j}}^{p_{ij}} (\mathbf{y}_i^a - \mathbf{y}_i^b)(\mathbf{y}_j^o - \mathbf{y}_j^b)/p_{ij},$$

Applied only along orbits in observation space

$\Delta t < 10$ min

2. Focus: a posteriori validation

Geometrical representation of error components

